Patterns of temperature anomalies at the Japan/East Sea surface, simulated under different kinds of wind forcing

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### Motivation

SSTA are usually considered as related to local air – sea heat exchange, however.... EOF analysis of the observed SST revealed that •there is the statistical relationship between SSTA and wind stress curl; •large-scale SSTA patterns in the Japan/East Sea can be related to dynamic processes – Ekman suction, front migrations, heat advection, etc...; •supported by qualitative considerations of simulation results from an oceanic

(Trusenkova et al., 2008).

What if we do EOF analysis of the simulated SST and compare with EOF modes of the observed SST?

model.

### Purpose of the study

- To simulate currents and tracers (temperature, salinity, etc ) in the Japan/East Sea, using an oceanic model forced by different patterns of wind stress curl.
- To compare EOF modes of the observed and simulated SST.
- To support suggested dynamic forcings of SST variability.

### MHI oceanic model

(Shapiro and Mikhaylova, 1992-1998)



 $b^{\sigma}_{k}$ ·is·base·buoyancy·of·the·kth·layer  $M_{0} = \alpha \Gamma - \beta R$  is surface buoyancy flux  $\Gamma$  is heat flux and R is freshwater flux

- 3D primitive equation, hydrostatic & Boussinesq
- Quasi-isopycnic co-ordinate in the vertical
- Complete thermodynamics, including
   surface heat/freshwater balances,
  - TKE model for the surface mixed layer,
  - prognostic equations for T and S,

- diapycnal exchange of T and S between the layers

- Variable T, S, and buoyancy in any layer
- Constraint on buoyancy variations in the inner layers (below the mixed layer)
- Bi-harmonic viscosity in the momentum equations
- Free surface
- Convective adjustment

### Simulation setup



- Lateral mesh of 1/8° (10-14 km) ~ marginally eddy resolving;
- 12 layers in the vertical;
- Bathymetry scaled by factor of 0.75;
- Depth of initial flat interfaces : 10, 25, 50, 75, 100, 150, 250, 350, 500, 700, 900 m;
- Initial T and S from average vertical profiles;
- Time step: 7.5 min;
- Bi-harmonic lateral viscosity: 2.5x10<sup>8</sup> m<sup>4</sup>/s, harmonic lateral diffusivity: 250 m<sup>2</sup>/s;
- Inflow transport in the Korean Strait after Takikawa and Yoon (2005), outflow in the Tsugaru and Soya Straits;
- Monthly atmospheric variables (1979-1999) for surface heat and freshwater balances;
- Wind stress from 1°-gridded NCEP/NCAR data;
- No restoration to the observed temperature or salinity.

### Simulation setup - 2

#### •Wind data:

NCEP/NCAR special product, SeaWiFS Project Ancillary Data, 6h, 1°x1°-gridded wind, 1998-2005,

34°-53°N, 127°-143°E (JES and adjacent land).

#### •Model runs:

- 11 years spin-up under the forcing of monthly wind 2000,

- 3 runs for 4 years under the forcing of different wind patterns but under the same surface buoyancy forcing and transports through the straits;
- results from 15th year of integration are analyzed.

#### Typical wind patterns (Trusenkova et al., 2007)



Warm

#### Monthly occurrence of the typical patterns



November – March and May → most frequent pattern April, June, July, October ??? → needs some additional considerations → analysis of seasonal variations of wind curl





The smallest scale  $\sim 40$  days

# Seasonal variation of wind stress curl

•Instantaneous occurrence: % of boxes over the JES, not including the adjacent land. A box adds to **C curl occurrence** if  $rot\tau > C_1$ ; **AC curl occurrence** if  $rot\tau < -C_1$ ; **weak wind occurrence** if  $-C_1 < rot\tau < C_1$ .

•Time series of occurrence for 1998-2005.

• Features:

2006

- opposition of C vs. AC curl occurrence (R = -0.67);
- semiannual periodicity;
- annual periodicity of the weak wind occurrence.

#### Run 0:

forcing by the semiannually varying wind stress curl



9

10

Wind patterns are timed at mid-months and linearly interpolated to every time step

# Run 1: forcing by the prevailing cyclonic curl

### Run 2: forcing by the westerlies



#### Run 0: monthly mean SSH

![](_page_11_Figure_1.jpeg)

#### Run 0: monthly mean SST

![](_page_12_Figure_1.jpeg)

#### Observed and simulated SST

Daily New Generation (NG) SST, Tohoku University, Sendai, Japan, July 1, 2002 - July 7, 2006, 34.5°- 48°N, 127.5°- 142°E, 0.05°-gridded, smoothed to 0.25° for computational purposes, satellite IR & MK

Daily Japan Meteorological Agency (JMA) **SST**, October 12, 1993 – November 8, 2006, 35°- 48°N, 127.5°- 142°E, 0.25°-gridded, satellite IR & MK + *in situ* 

similar patterns from both datasets (Trusenkova et al., 2008)  $\rightarrow$  JMA SST compared to the simulated SST

Daily SST in the 15th year of integration from Runs 0 - 2.

#### Decomposition to EOF:

#### $X(r, t) = \sum A_k(r) \cdot B_k(t),$

where r stands for the spatial coordinates and t for time, X(r, t) is the original SST.

•Correlations, not covariances  $\rightarrow$  detection of low amplitude SSTA

in the northern JES.

•Residual anomalies:  $X_a(r, t) = X(r, t) - A_1(r) \cdot B_1(t)$ ,

![](_page_14_Figure_6.jpeg)

![](_page_14_Figure_7.jpeg)

•Anomaly related to the kth mode:  $X_r(t) = A_k(r_0) \cdot B_k(t)$ , where  $r_0$  is characteristic location, usually within the spatial core.

#### Data processing

## EOF of JMA and simulated SST:

![](_page_15_Figure_1.jpeg)

- Modes 2 and 3 are degenerate in Run 0.Modes 2 and 3 are switched in Runs 0 & 1.
- •Statistically significant Modes 1 3 of JMA SST.

Compare spatial patterns:

•Mode 1 of JMA vs.

Mode 1 of Runs 0 - 2.

•Mode 2 of JMA vs. Mode 3 of Run 0.

•Mode 3 of JMA vs. Mode 2 of Runs 0 & 1.

•Mode 3 of Run 1 and Modes 2 & 3 of Run 2 are different from JMA and Run 0.

•Simulated SST: interannual variability is absent, mesoscale variability is low.

•High fraction of variance accounted for by Mode 1 of the simulated SST.

•Spatial cores are diffused for JMA SST and are stretched out along currents/fronts for the simulated SST.

![](_page_16_Figure_0.jpeg)

#### Mode 1: forcing by Ekman suction/pumping

![](_page_17_Figure_1.jpeg)

•SSTA is largest in May-June & November-December and smallest in February-March & August-September.

•Northward shift of Subarctic Front in spring (Nikitin, 2002); formation of the Northwestern Branch in autumn (Danchenkov et al., 1997).

•Western core around Subarctic Front in Run 0 is forced by the AC curl in spring and autumn.

•No western core in Run 1 under the forcing of the prevailing C wind curl.

•Spatial/temporal patterns are

# Mode SAF – North: forcing by AC wind stress curl

Occurrence of the AC & C wind curl in the western JES

![](_page_18_Figure_7.jpeg)

![](_page_19_Picture_0.jpeg)

northarnmost IFS

# Northern vs. Western Cores: opposition in wind curl and SSTA

Occurrence of the AC/C wind stress curl AC curl C curl Western JES •Opposition of wind stress curl over the central and northernmost JES. Northern JES •Semiannual wind stress curl in the northernmost JES in Runs 0 & 1. •Variability around the eastern 17.6 % 10.8% 16.4% 9% margin of the cyclonic gyre in EOF2 EOF2 EOF3 EOF3 Runs 0 & 1. •Spatial/temporal patterns JMA Run 2 Run 0 Run 1 are different in Run 2. SSH JMA September •AC eddy formation (Nikitin, Run 0 Run 1 2007) and variety of circulation Run 2 patterns (Dyakov, 2006) in the

Pathways of warm water transport to the Russian coast

![](_page_20_Figure_1.jpeg)

(Lobanov et al., 2001)

Forcing by the prevailing C wind stress curl in Runs 0 & 1 but not in Run 2

Month

().1 L<sup>2</sup> JMA Run 0

Run 1 Run 2

1993

![](_page_20_Figure_4.jpeg)

200

### Conclusion

•EOF modes of SST simulated under the forcing of realistic wind are consistent with those of the observed SST.

•Northward shift and a secondary branch of Subarctic Front in the western JES are simulated under the forcing of the local AC wind stress curl. These processes result in formation of positive SSTA.

•Semiannual periodicity in SST is simulated under the forcing of the semiannually varying wind stress curl.

•Northward transport of warm water along the Honshu coast is increased and accompanied by the decreased transport in the western JES in late summer – autumn under the forcing of easterly wind patterns with the prevailing C curl.

•To estimate a contribution of mesoscale dynamics to the large-scale SST modes, truly eddy-resolving simulations are necessary.

•Interannual variability of the observed SST can be related to the interannual wind variations.

# Wavelet analysis of the occurrence of wind curl

![](_page_22_Figure_1.jpeg)

8485181524

Power

![](_page_22_Figure_2.jpeg)

#### **Residual anomalies SSTA:** JMA vs. NG Residual anomalies calculated by subtraction of an average annual cycle: 40.6% $\mathbf{X}_{a}^{(1)}(\mathbf{r}, t) = \mathbf{X}(\mathbf{r}, t) - A_{1}^{(0)}(\mathbf{r})B_{1}^{(0)}(t)$ Temperature °C °C MODE 1 1.4 EOF 1 NG South 2 JMA South 0 17.6 % 0.6 NG NW -2 A Property Loss JMA NW 0.2 -0.2 MODE 2 -0.6 EOF 2 2 -1 6.2 % 0 -1.4 -2 NG North West G -6 JMA West JMA North MODE 3 2 EOF 3 -2 NG Honshu JMA Honshu

2005

2006

2002

2003

2004

### JMA SSTA: correlations vs. covariances

![](_page_24_Figure_1.jpeg)

Residual anomalies calculated by subtraction of an average annual cycle:  $\mathbf{X}_{\mathbf{a}}^{(1)}(\mathbf{r}, t) = \mathbf{X}(\mathbf{r}, t) - A_1^{(0)}(\mathbf{r})B_1^{(0)}(t)$ 

![](_page_25_Figure_0.jpeg)

#### SSH: Run 0 vs. Run

![](_page_26_Figure_1.jpeg)

#### **Run 1**:

No seasonal shift of Subarctic Front

Subarctic region extends far south in the western Japan Sea